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Lithofacies And Geochemical Properties Of Neogen Deposits At South Of Tuzgolu- Turkey

Arif Delikan¹, Muazzez Celik Karakaya¹, Necati Karakaya¹, Hatice Unal Ercan¹, Ayla Bozdogan¹

Abstract

The Tuzgolu Basin located at the Central Anatolia (Turkey) is bounded by Ankara uplift at the north, the Kirsehir massif from the east and the Sivrihisar-Bozdogan massive at the west. In the study area which is located at the South of the Tuzgolu, the Paleogene and Mesozoic marine carbonates and igneous rocks underlies the Neogene sequences. Neogene deposits consist of Kizilbayir, Katrandetepe and Bestepeler formations which are conformable with each other. 10 different lithofacies were identified within the Neogene sequence by considering sedimentation conditions, lithology, sedimentary structure and fossil content.; Grain-supported conglomerate facies (Gcu), Convolute bedded sandstone facies (Sk), Thick-bedded sandstone facies (St), Gray-purple colored thick layered mudstone facies (Mt), Oolitic limestone facies (OC), Alternating gypsum-anhydrite-mudstone-micritic limestone facies (Cmag), Bituminous shale facies (Bs), Halite-mudstone facies (Hm), Massive and parallel laminated tuff facies (Pmlt), Alternating mudstone-sandstone facies (Ms). The facies analysis show that sedimentation in the study area began with fluvial sediments (Kizilbayir formation) and followed by sediments of shallow lake which was often interrupted by sediments from land (Katrandetepe formation), and by the interbedded mudstone, sandstone, conglomerates and tuff at the closure of the lake (Bestepeler formation). According to the geochemical analysis results obtained from lake carbonate and evaporite deposits (Halite, anhydrite and gypsum), REE, LILE and HFSE values are more abundant in clayey samples than those in other evaporitic sediments. The Sr contents of halites (1-1539 ppm) are lower than sulfate (183-4378.04 ppm) and carbonates (922-12365 ppm). Halite minerals contain very high Cl (505686-615905 ppm) and low Br (5-637 ppm) indicating that they are products of dissolution, mixing and re-precipitation.

Keywords: Tuzgolu, Lithofacies, Halite, Gypsum, Anhydrite

1. INTRODUCTION

The Anatolian plateau was formed by the collision of the Arabian and Eurasian plates ([2] and [3]). The study area is located at the Central Anatolia is an inner enclosed basin and is bounded by Ankara uplift in the north, the Kirsehir massif from in the east and the Menderes massif from the west (Figure.1). During this time, two major fault systems in the area, the Tuz Golu and the Sultanhanı faults, developed as south-west dipping, NW–SE striking, normal faults. At some time in the Late Miocene-Early Pliocene, during regional subsidence, a previously unreported phase of contraction occurred, which led to the development of a north-east–vergent thrust sheet, the culmination of which forms the morphologic ridge to the east of the Tuz Golu Lake ([2], [3], [4], [5], [6], [7], [8] and [9]).

The Tuz Golu Basin is located in the south-eastern part of Central Anatolia (Figure 1). Evaporitic and carbonate (limestone and dolomite) deposits are deposited from Na, Ca, Mg, Cl, CO₃, and SO₄ in different contents and most of them were precipitated inner continental basin, tectonically active (slump folding, chaotic structure and collapse deposits) and arid lacustrine environments which effected time to time intake of seawater ([1], [2], [3]). In this study area, more thicknesses of salt and soda than a few hundred meters were deposited during Miocene period.

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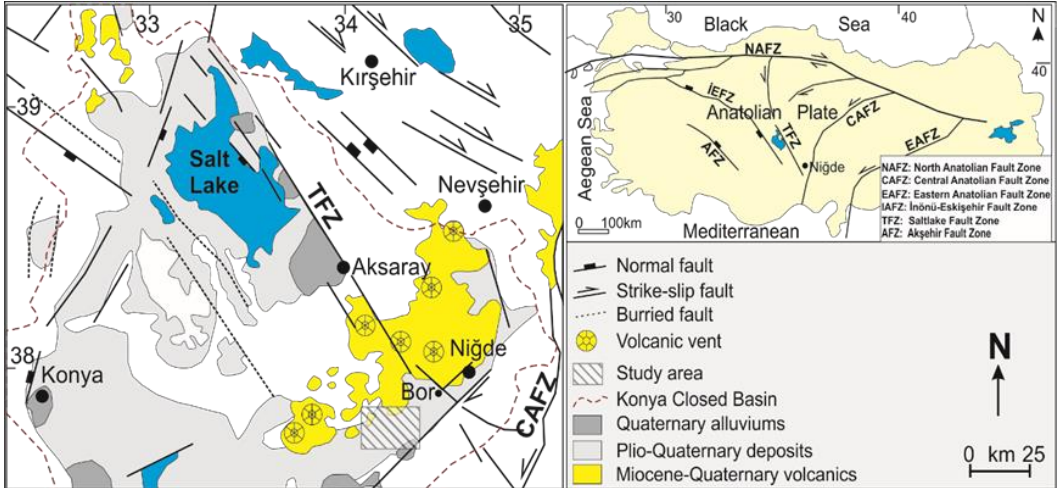


Figure 1. Simplified geology map (after [1])

2. RESULT AND DISCUSSION

2.1 STRATIGRAPHY

There are sediments of more than 3000 meters deposited in the Paleozoic-Quaternary period at the base of the Tuz Golu basin and various rock assemblages settled with volcanic activity (Figure 1).

2.2 Basement Rocks

Paleocene (Serenkaya formation), Eocene (guney formation) and Paleozoic aged metamorphic and sedimentary rocks (Asagidedigi formation) are found at the base of the late Miocene aged Lake formations (Figure 2 and 3)

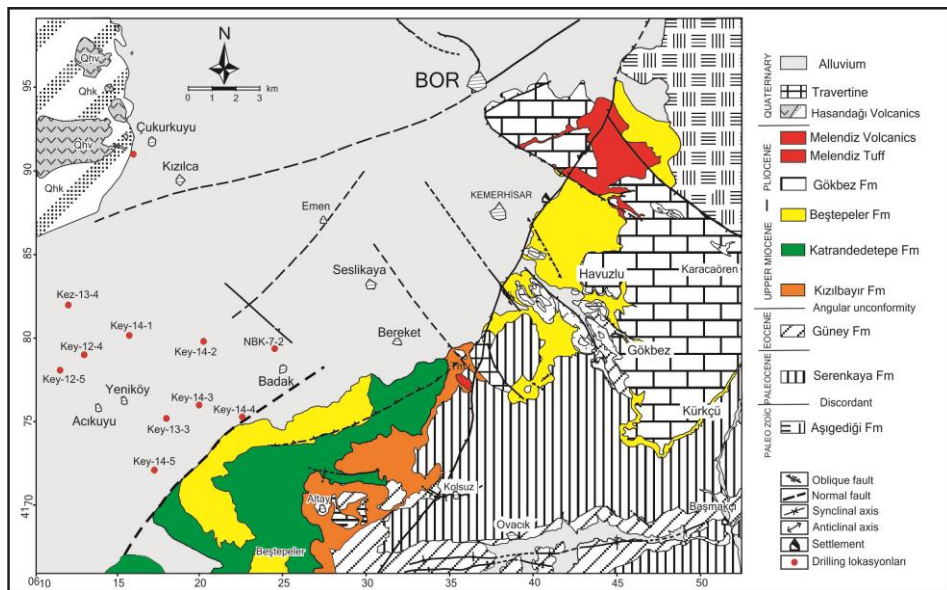


Figure 2 Geological map of study area (from [4], [5], [6], [7],) and drill locations

ERATHM	SYSTEM	SERIES	STAGE	FORMATION	THICKNESS	LITHOLOGY	EXPLANATION	
CENOZOICK	NEOGENE	QUATERNARY			20-30		Gravel-sand-mud accumulating in valley bottoms	
					15-20		Poorly cemented gravel and sand	
		PLIOCENE	HASANDAG VOLCANICS				Travertine	
							Pyroclastic-Basalt and andesites	
			MELENDEZ VOLCANICS				Basalt and andesites	
							Tuff and agglomera of Melendez volcanics	
			GÖKBEZ				Abundant plant fossiliferous mudstone-marl alternation	
							Tuff and aglomera of Melendez volcanics	
			BESTEPELER		400-450		Green-gray colored sandstone with gypsum and anthracite bands Mudstone alternation	
							Polygenic conglomerate and sandstone with large gravels alternations	
			KATRANDEDETEPE				Beige-Green lacustrine limestone marl-mudstone alternations	
					700-1400?		Gypsum and mudstones starting with the basal conglomerate	
						Anhydrite and mudstone alternation with coal level.		
						Glauberite-anhydrite and halite-mudstone alternation		
		KIZILBAYIR		200-350		Bituminous shale with dark brown oil spill. Dolomite and clayey limestone		
						Carbonized level with a thickness of about 30 cm		
		Eocene	MIDDLE	GÜNEY	600-1000		Reddish green colored claystone-siltstone alternation	
							Green colored, cross-bedded and laminated sandstones	
		PALEOSEN	UPPER	SERENKAYA	800-1200		Well cemened poligenic conglomerate	
							Gray colored mudstone, thick bedded sandstone and shale alternations	
MESOZOIC	CRETACEOUS	SİNEKSİZYAVLA METAGABRO			Sedimentay facies with limestone blocks			
					Conglomerate-sandstone with normal grading and Shale alternation			
PALEZOIC		AŞICEDİĞİ			Coarse-grained conglomerate contain magmatic rock pebbles			
					Metagabro			
						Crystalline limestone		

Figure 3. Generalized stratigraphic section of the study area (from [4], [5], [6], [7]).

2.3 Kizilbayir formation

The formation comprises reddish and yellow colored conglomerate, sandstone and mudstone. Additionally, in the drilling logs, normal grading of the conglomerates, cross bedding and lamination in the sandstones are observed.

The Kizilbayir formation was deposited in a tectonically very active basin so that depositional conditions changed frequently and several different facies were deposited.

Three lithofacies were defined within the Kizilbayir Formation;

1. Clast-supported conglomerate lithofacies (Figure 4)
2. Convolute bedding sandstone lithofacies (Figure 4)
3. Gray-claret colored thick-bedded mudstone lithofacies (Figure 4)

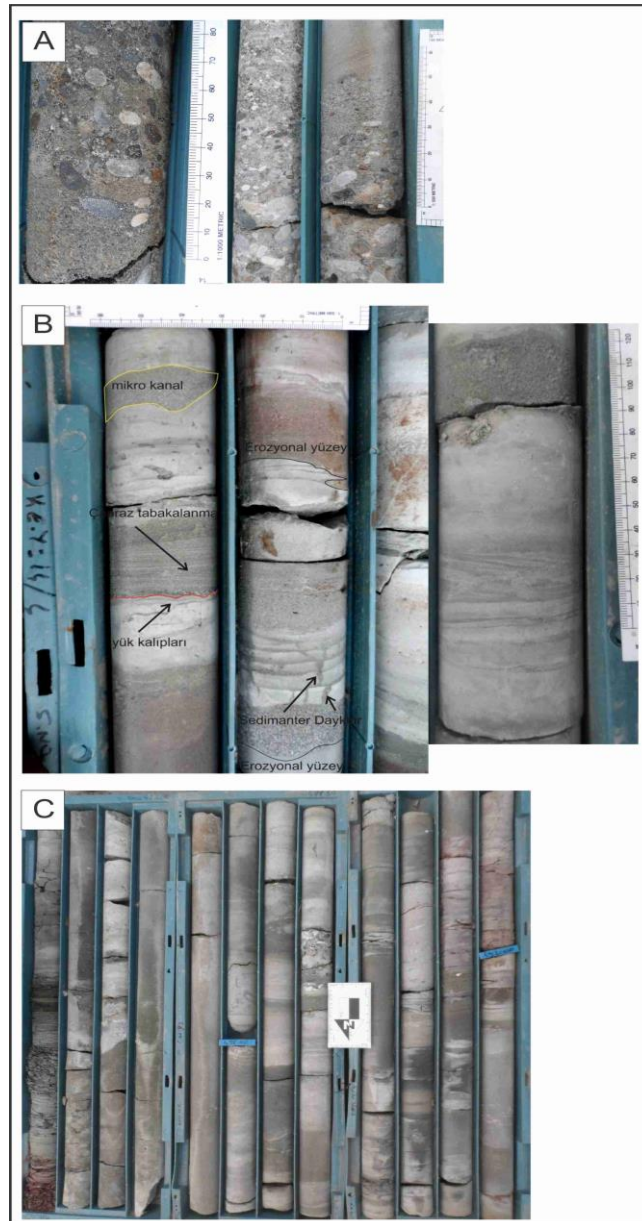


Figure 4 Lithofacies of Kizilbayir formation; A. Clast-supported conglomerate lithofacies (Normal graded conglomerate), B. Convolute bedding sandstone lithofacies, C. Gray-claret colored thick-bedded mudstone lithofacies

2.4 Katrandedetepe formation

Formation consists of halite, gypsum, anhydrite, gulberite, mudstone, limestone, dolomite, bituminous, petroliferous sandstone, siltstone and sandstone. The formation commonly contains deformational structures. Such as salt diapirs, slump folds, and load cast structures. Three lithofacies were defined within the Katrandedetepe formation;

1. Oolitic limestone lithofacies (Figure 5)
2. Anhydrite, gypsum –mudstone –micritic limestone lithofacies (Figure 5)
3. Bituminous, petroliferous sandstone (Figure 5)

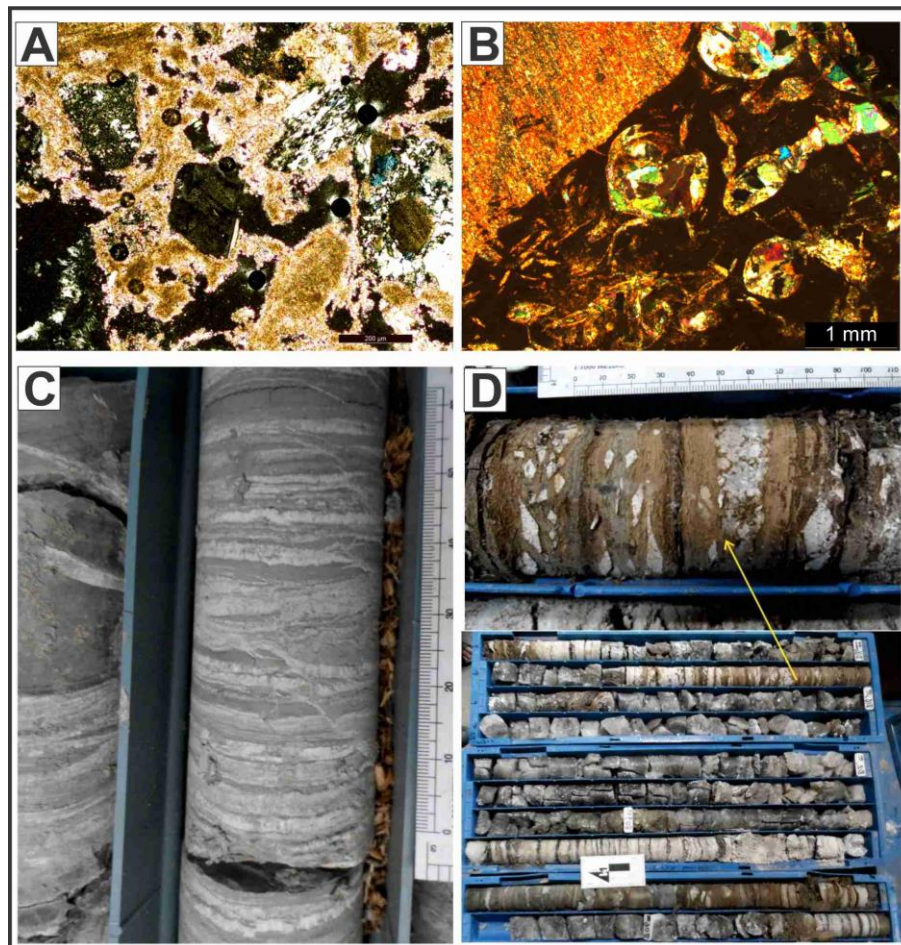


Figure 5 Lithofacies of Katrandedetepe formation; A. Oolitic limestone lithofacies (Thin section photos), B. Anhydrite, gypsum –mudstone –micritic limestone lithofacies (Gastropoda shell was filled by anhydrite), C. Bituminous, petroliferous sandstone Lithofacies

2.5 Bestepeler formation

The formation includes light green mudstone, siltstone, sandstone, conglomerate, cream colored tuff and tuffite.

Three lithofacies were defined within the Bestepeler formation;

1. Halite-mudstone-lithofacies (Figure 6)
2. Masiv and tabular bedding lapilli tuff lithofacies (Figure 6)
3. Mudstone-sandstone alternate lithofaciess (Figure 6)

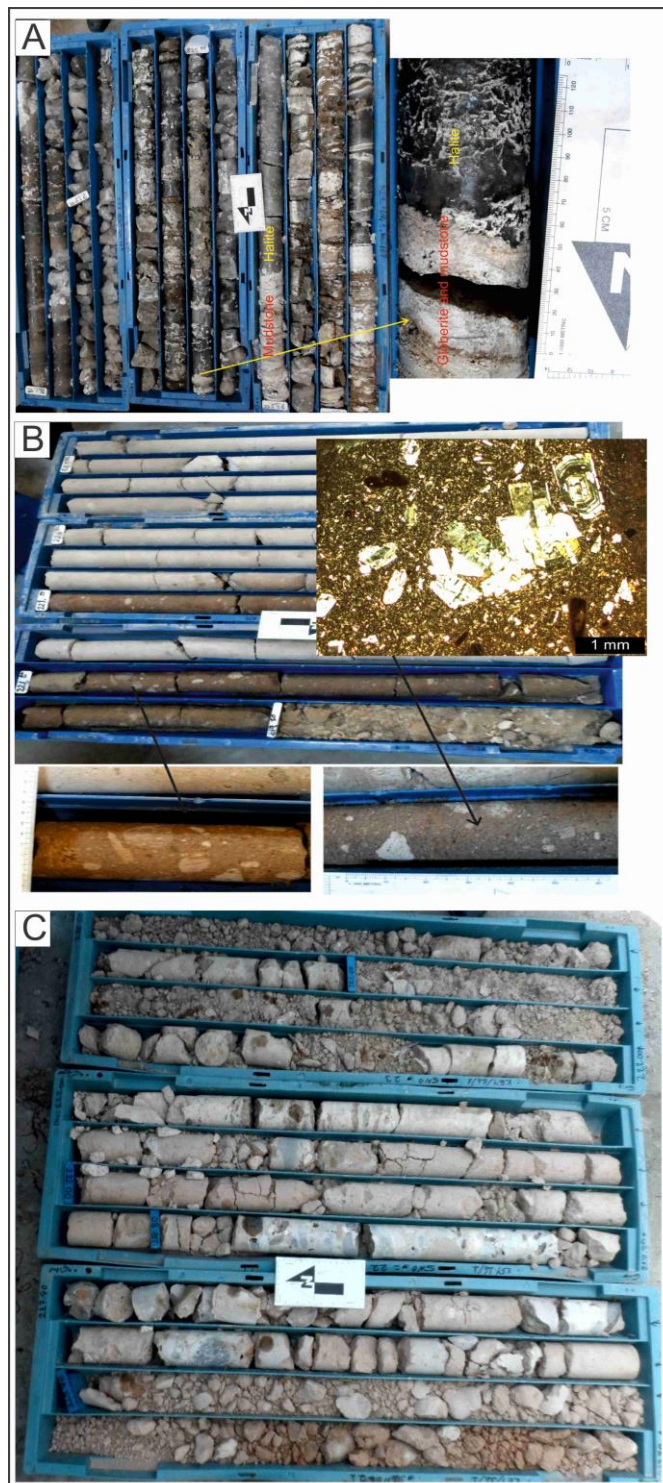


Figure 6 Lithofacies of Bestepeler formation; A. Halite-mudstone-lithofacies, B. Masiv and tabular bedding lapilli tuff lithofacies, C. Mudstone-sandstone alternate lithofaciess

3. GEOCHEMICAL PROPERTIES

According to petrographic and geochemical analyzes, 3 groups of minerals were identified. distribution of minerals according to depths in drilling log (Figure 7). Three groups are Sulphates, carbonates and Halite-Globerite

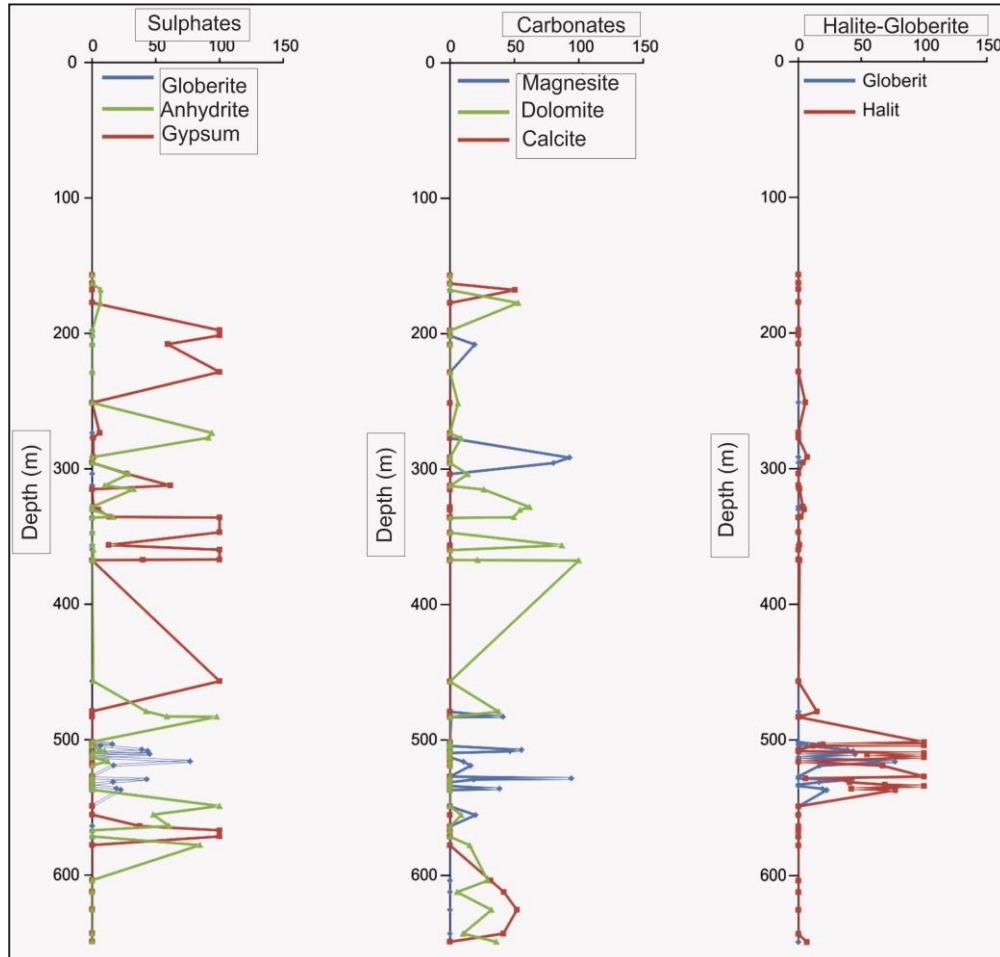


Figure 7. Changes of mineral contents of TG3 drilling.

4. CONCLUSION

1. The regional subsidence phase in the study area started in Tortonian times. A sediment sequences (Bestepe, Katrandetepe and Kizilbayir formations) which is more than 1000 m thick deposited from Late Miocene and possibly into Pliocene times
2. Facies properties and the synsedimentary structures (such as slump folding, chaotic levels, sedimentary dykes ect.) in the Katrandetepe formation show that the region was tectonically controlled.
3. The sequences comprise alternating evaporitic deposits (the cycle of the halite, gypsum, anhydrite) and siliciclastic deposits. This indicate that there was a continuous sedimentation which was mainly controlled by paleoclimatic changes

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